

RESEARCH PROCESS



BL KINEMATIC VARIABLES

- Gust Factor
 - Also used to convert between different averaging times → “Durst Curve”

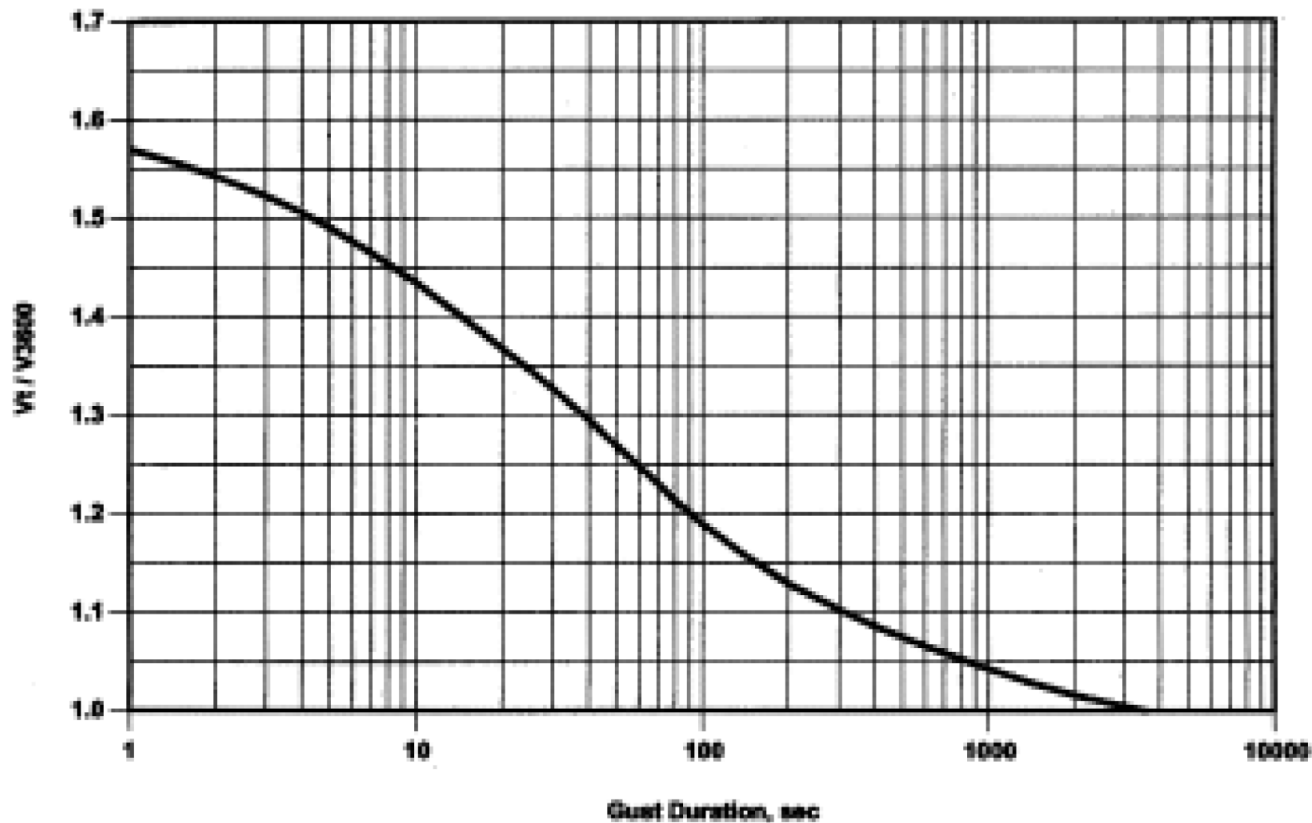


FIGURE C6-4 MAXIMUM SPEED AVERAGED OVER t_g TO HOURLY MEANS SPEED



THE RESEARCH PROCESS

- Develop research question (s)
 - “Science Leads.”
 - For instance, “ Why do some thunderstorms produce tornadoes, while others do not?”
- Narrow down your questions and develop *TESTABLE* hypotheses
 - Example: “ Thunderstorms need strong low level winds to produce a tornado.”
- Determine what data you need and how you will collect it
 - The data will dictate what tools you need. Our focus is on wind, so what tools allow us to collect wind data?



THE RESEARCH PROCESS

- You know what tools you need and what kind of data you will collect, but how will you collect it?
 - Field Projects
 - Example: Project SCOUT
- Data collection must be carefully designed and must take into account personnel and equipment safety.
- Important:
 - Instrument characteristics
 - Data
 - Biases



THE RESEARCH PROCESS

- Data collection is over...now what?
- DATA PROCESSING
- The ultimate goal of research is to share your answers with the scientific community
 - Publications
 - Conference Presentations
 - Lectures

ORIGINAL RESEARCH article

Front. Built Environ., 15 October 2019 | <https://doi.org/10.3389/fbuil.2019.00119>



Exploring the Feasibility of Using Commercially Available Vertically Pointing Wind Profiling Lidars to Acquire Thunderstorm Wind Profiles

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1438

MONTHLY WEATHER REVIEW

VOLUME 109

Five Scales of Airflow Associated with a Series of Downbursts on 16 July 1980

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(Manuscript received 5 September 1980, in final form 17 March 1981)

ABSTRACT

A series of destructive windstorms on 16 July 1980 in a 50 km (30 mi) wide zone from Chicago to Detroit was surveyed both from the air and the ground. In spite of the initial suspicion of 10–20 tornadoes in the area, the nature of the windstorms was confirmed to be downbursts and microbursts characterized by multiple scales of airflows with their horizontal dimensions extending tens of meters to hundreds of kilometers.

An attempt was made to estimate the wind speed based on three types of airborne objects: a 180 kg (390 lb) chimney, a 1000 kg (one ton) corn storage bin, and lumber from damaged roofs found inside downburst areas, obtaining the maximum wind speed of $63 \pm 10 \text{ m s}^{-1}$ ($140 \pm 25 \text{ mph}$). A total of \$500 million damage reported was caused by thunderstorm-induced non-tornadic storms which affected very large areas.

SMS/GOES pictures showed that the parent cloud was oval-shaped with its lifetime in excess of 12 h. The overshooting areas enclosed by the -66°C isotherms shrunk rapidly at the onset of the Chicago-area downbursts, indicating that the downbursts began when overshooting activities subsided. This variation of the overshooting features, however, does not necessarily imply a direct physical link between the collapsing top and the downbursts at the surface. This paper presents cloud-top features and wind effects on the ground with no attempt to relate them on the basis of conceptual models currently available.

1. Introduction

Suckstorf (1938) postulated that precipitation cooling within thunderstorms causes an outflow of cold air which results in strong surface winds. As air traffic increased in the late 1930's, squall-line related accidents occurred in various parts of the world, resulting in the operation of fact-finding projects of thunderstorms and squall-line circulations. Results of the Japanese Thunderstorm Observa-

one related the localized wind with the foot of a strong downdraft which could spread out violently and result in uprooted trees and damaged houses.

There are numerous reports of uprooted trees in NOAA's *Storm Data* in which types of damaging winds are classified as tornado or straight-line winds. Straight-line winds are assumed to be those which rush out of thunderstorms behind advancing gust fronts which are tens of kilometers in length.

EARTH'S CLIMATE HISTORY (AND PRESENT)



EARTH'S CLIMATE HISTORY

- Based on dating of Uranium isotopes, Earth has been around for about 4.6 billion years.
- Most of the paleo-climatological methods (proxies) only give us information from ~ the last million years.
 - Dendroclimatology / Dendrochronology
 - Ice Cores



GOOD QUOTE

- Data from before the instrument record are called proxy data.
- Proxy Data – “ *...long-lived geological, chemical, or biological systems that have the climate imprinted on them*” -- Dessler (2016)
- Goal: Reconstruct past climates and analyze the results

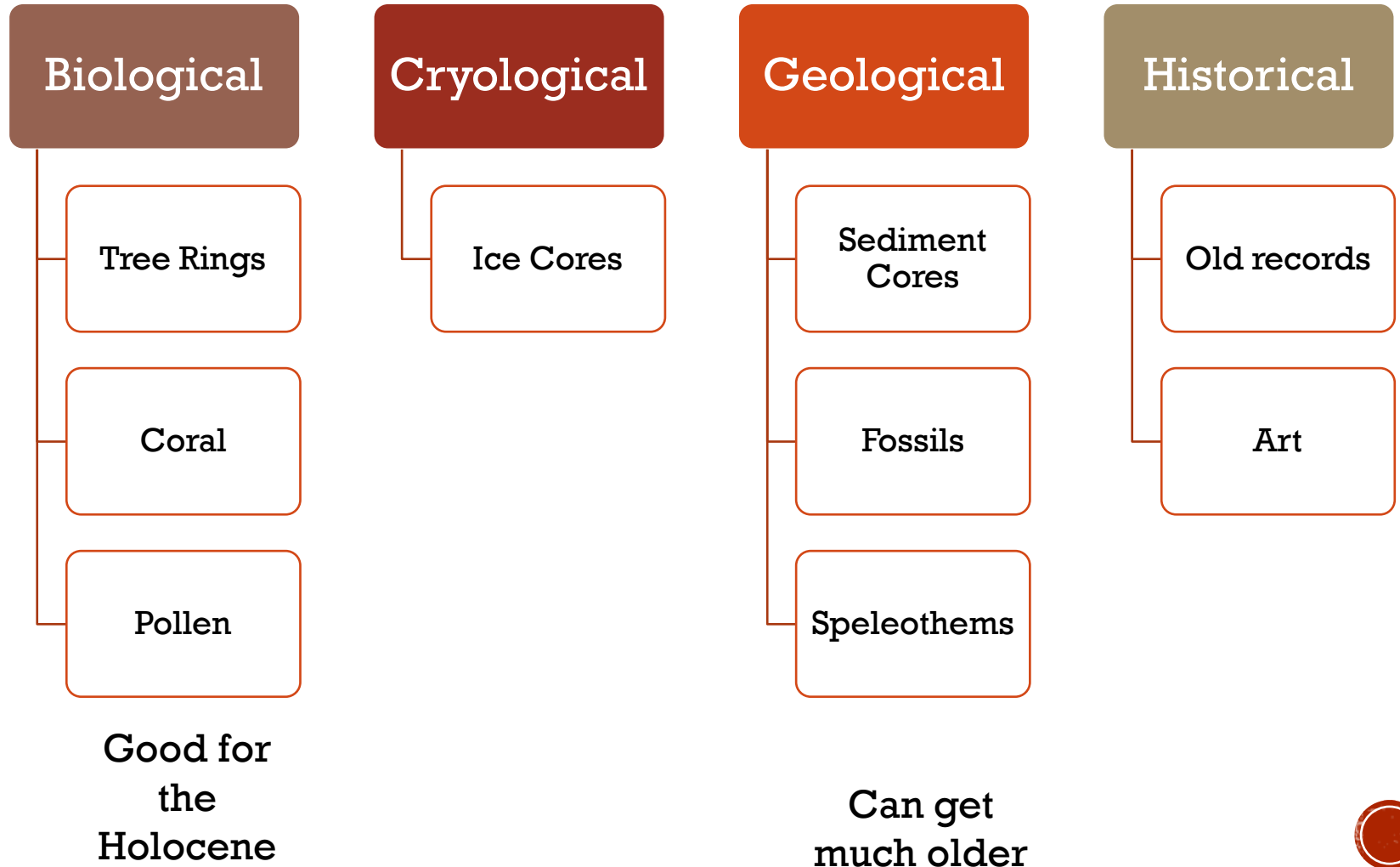


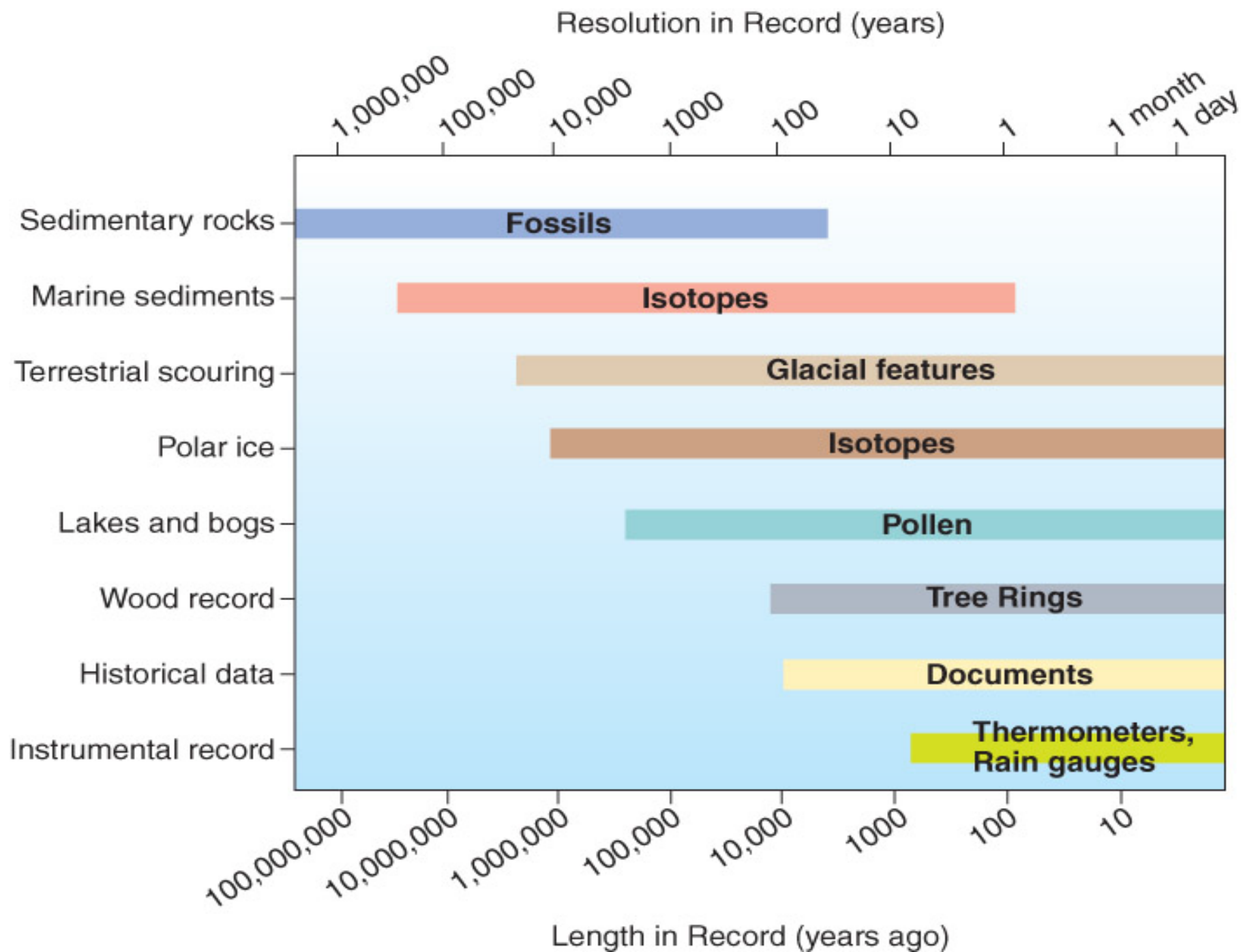
PALEOCLIMATOLOGY

- Proxy Data:
 - Accuracy declines with increasing time before present.
 - Declining Resolution
 - Fragmented records
 - Correlating proxy variables to physical quantities.
 - Lack of benchmarks



PALEO-CLIMATOLOGY





DENDRO-

- -chronology: the use of tree rings for dating.
- -climatology: the use of tree rings as a proxy indicator of climate.
- Both based on analyzing the thickness and density of annual growth rings of certain tree species.

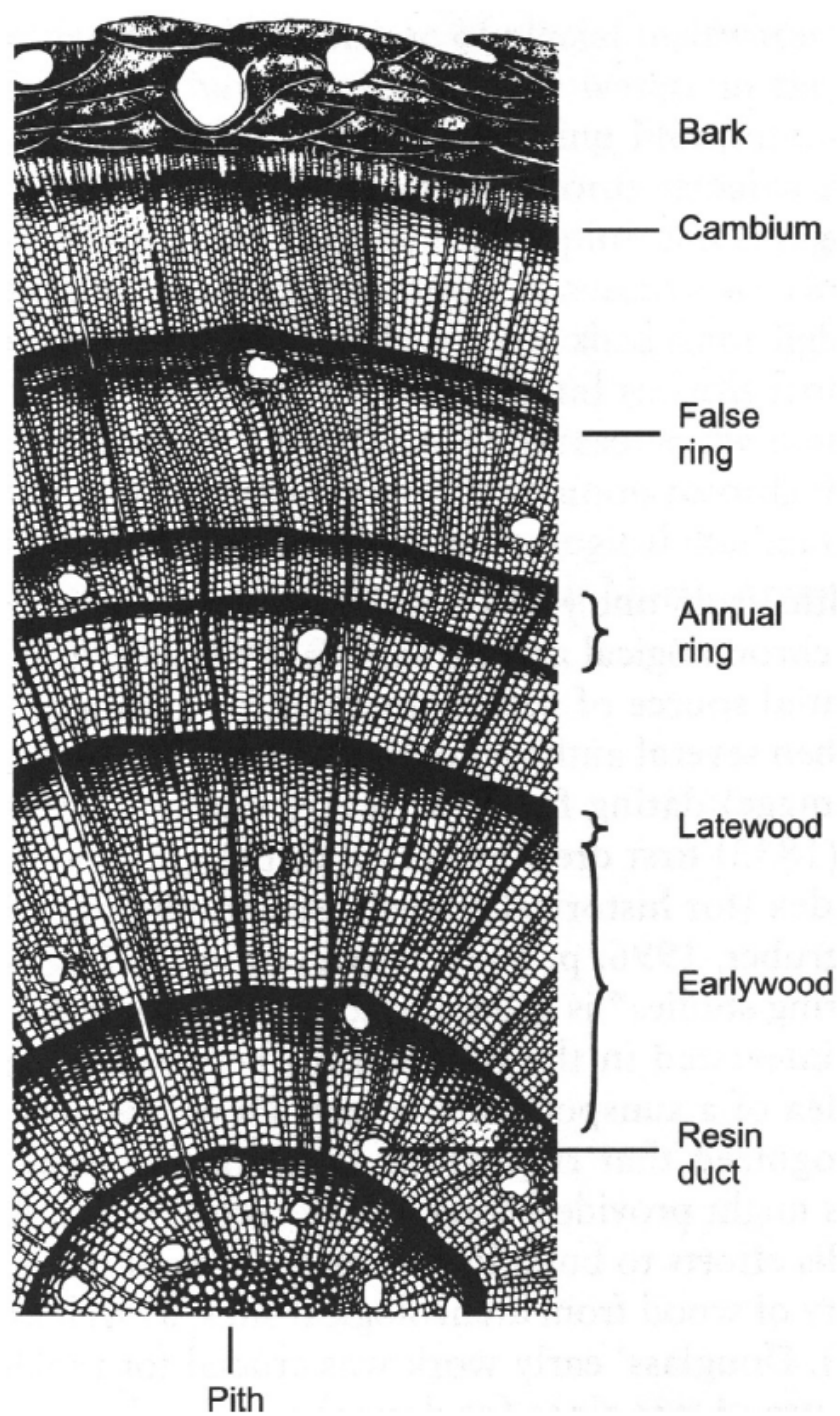


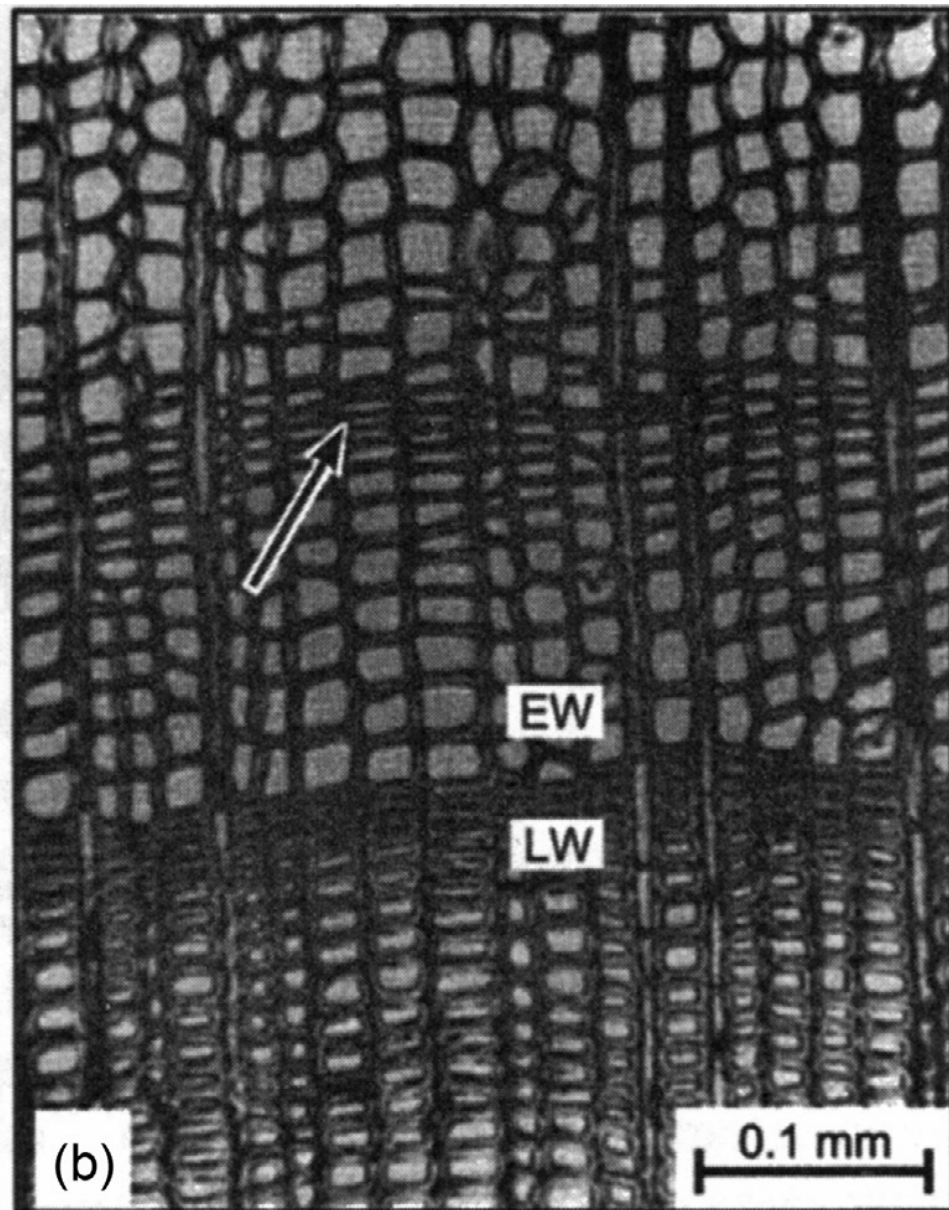
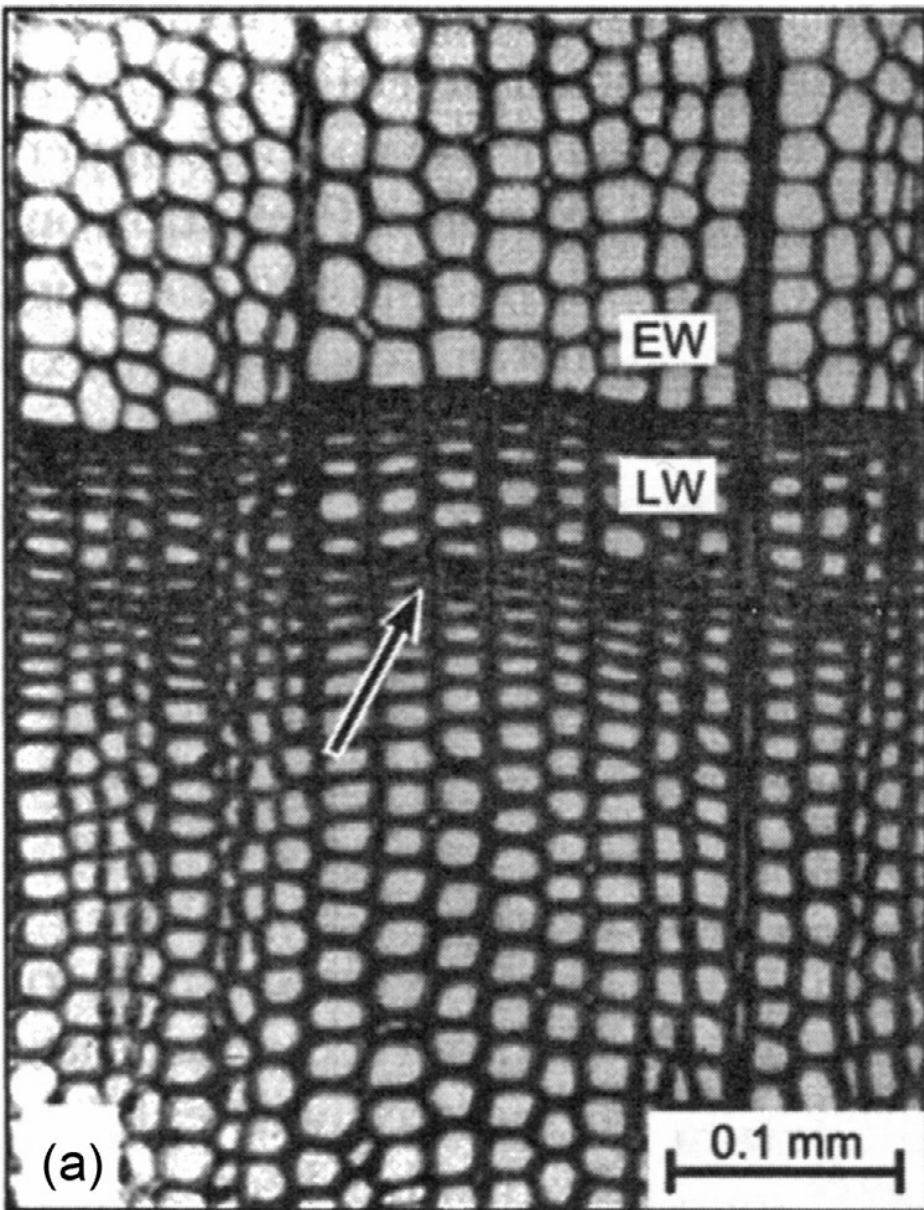


TREE RINGS: LIGHT AND DARK

- Color based on cell type:
 - Spring growth cells are lighter in color
 - Late growth cells are smaller, thick-walled, and more dense
- Abrupt changes in cell type mark the boundary between annual growth rings.
- Can get “false rings” when conditions get rough before end of growing season.







TREE RINGS

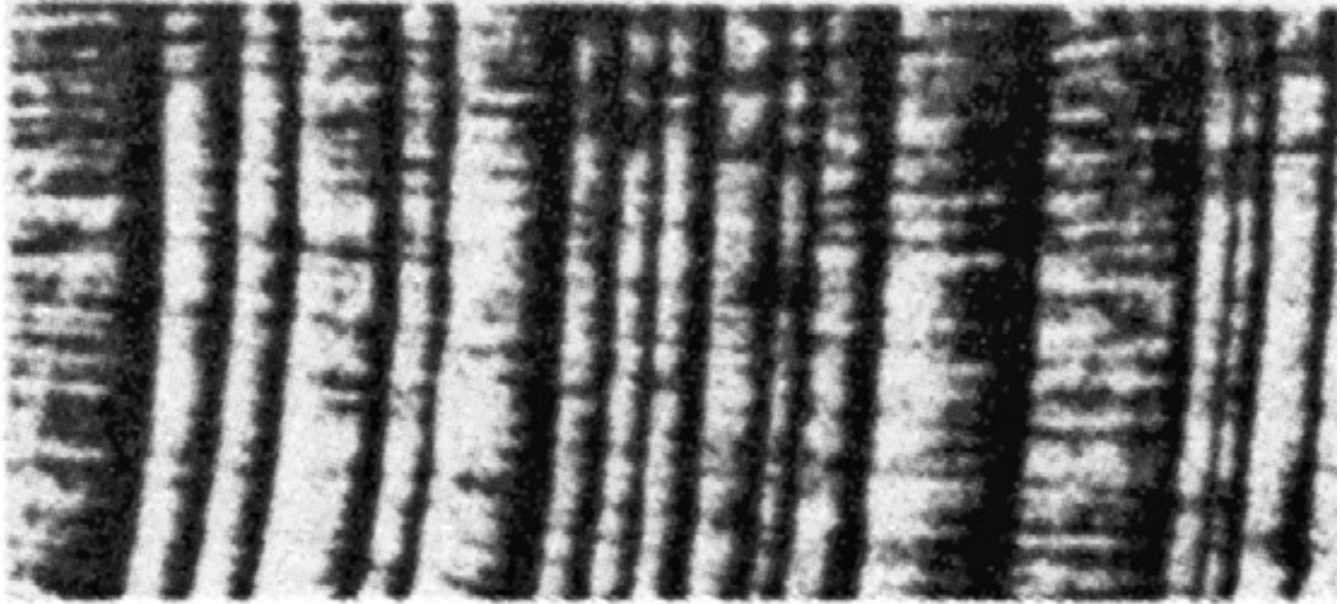
- The width of the rings is controlled by multiple factors:
 - Age of the tree
 - Tree species
 - Food storage / soil quality
 - Climate factors
 - Sunshine Temperature
 - Precipitation Humidity
 - Wind speed
- Account for non-climate signals through the tree-growth index: ratio of width to expected width based on age



ISOLATING A CLIMATE SIGNAL: SITE CONSIDERATIONS

- Need a location where trees are under some stress:
 - Temp stress: trees @ altitudinal or latitudinal tree line
 - Precipitation stress: trees in semi-arid regions
- Example: Proximity to the water table can reduce the climate signal.





ISOLATING A CLIMATE SIGNAL: SITE CONSIDERATIONS

- Need a location where trees are under some stress:
 - Temp stress: trees @ altitudinal or latitudinal tree line
 - Precipitation stress: trees in semi-arid regions
- Example: Proximity to the water table can reduce the climate signal.
- Location with a proper sample size:
 - 2 to 3 cores from each tree
 - ~20 trees from each site.



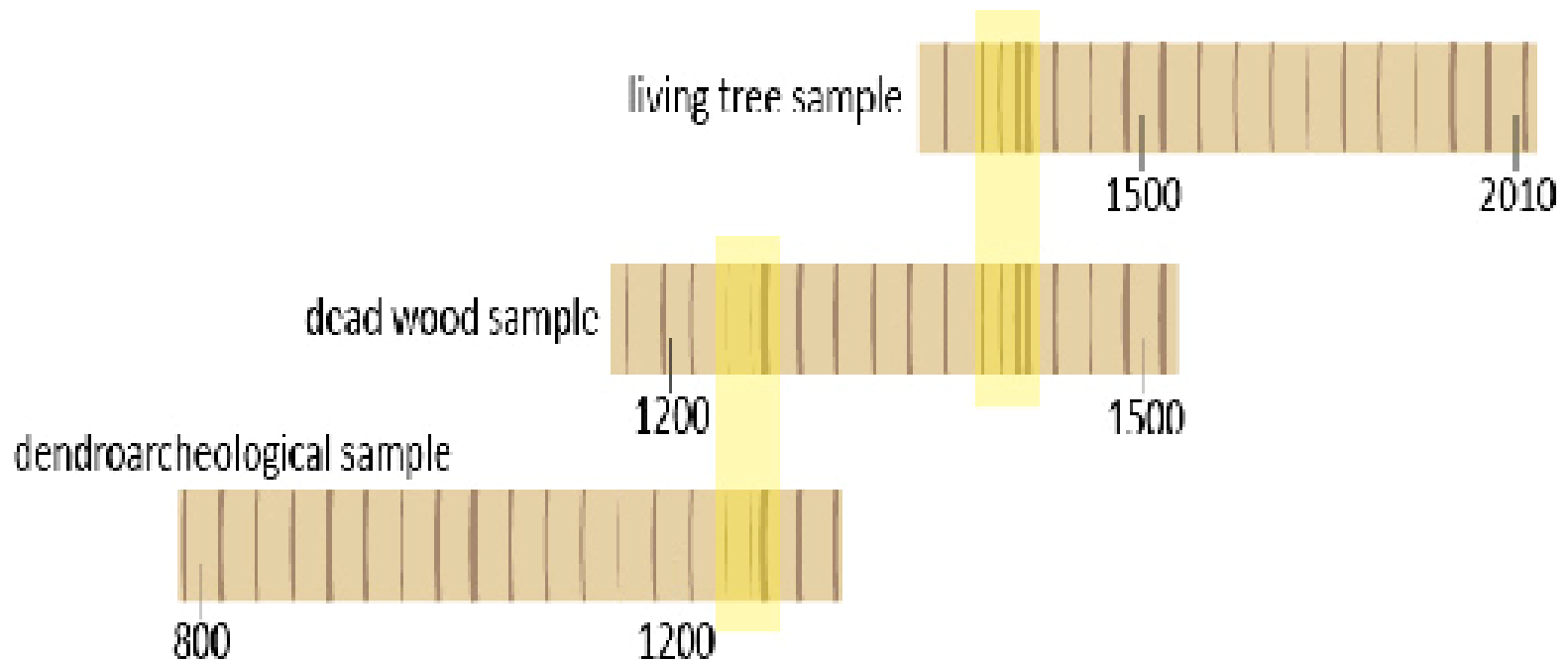
ISOLATING A CLIMATE SIGNAL: SITE CONSIDERATIONS

- Primary Tree species used are:
 - Ponderosa Pine
 - Douglas Fir
 - Bristlecone Pine
- Oldest living individual tree is ~ 5,000 years old.



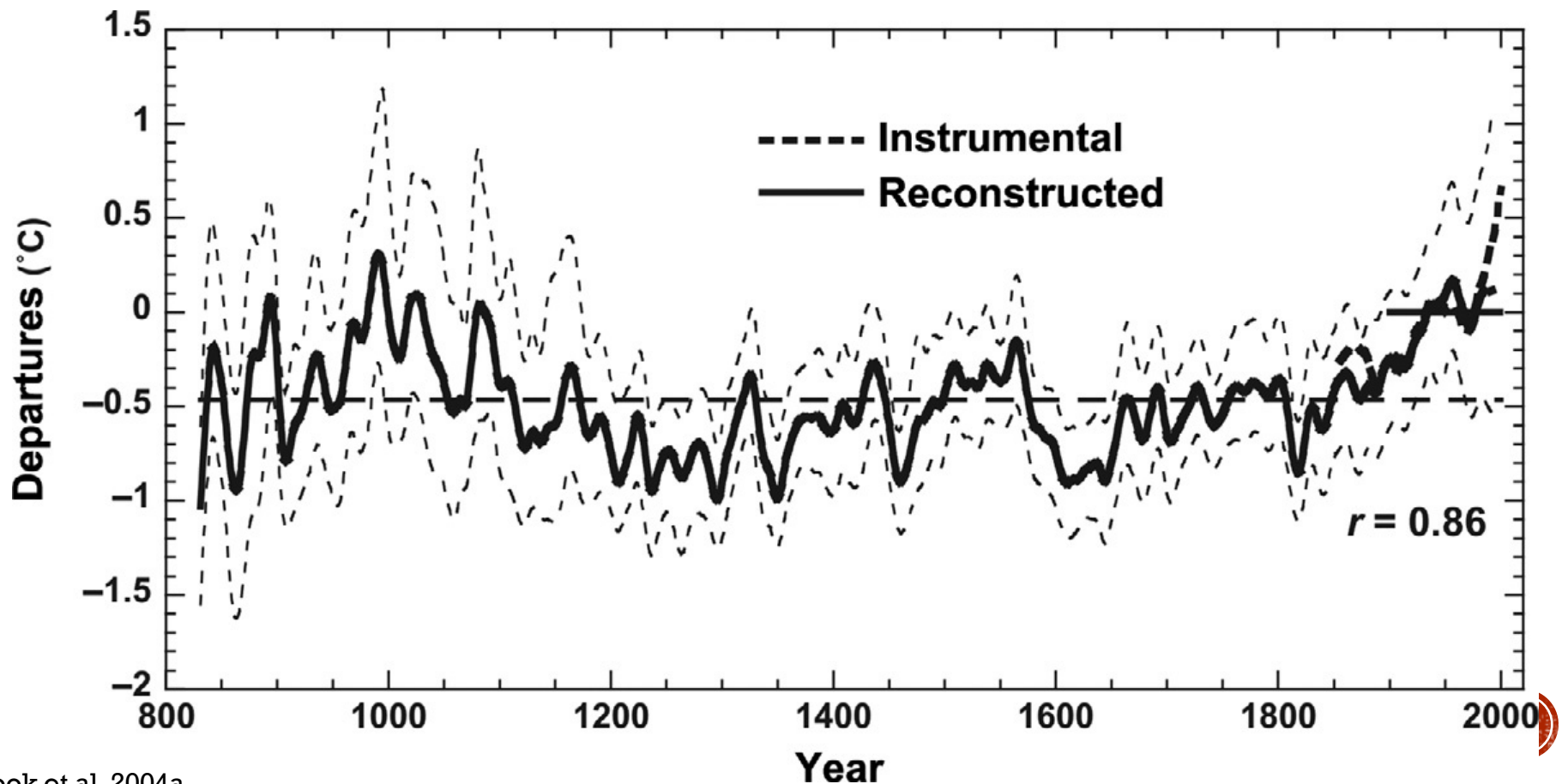
TREE CORE RECORDS

- Records are typically limited to 500-700 years.
- Can use “cross dating” to extend the record back ~10,000 years.



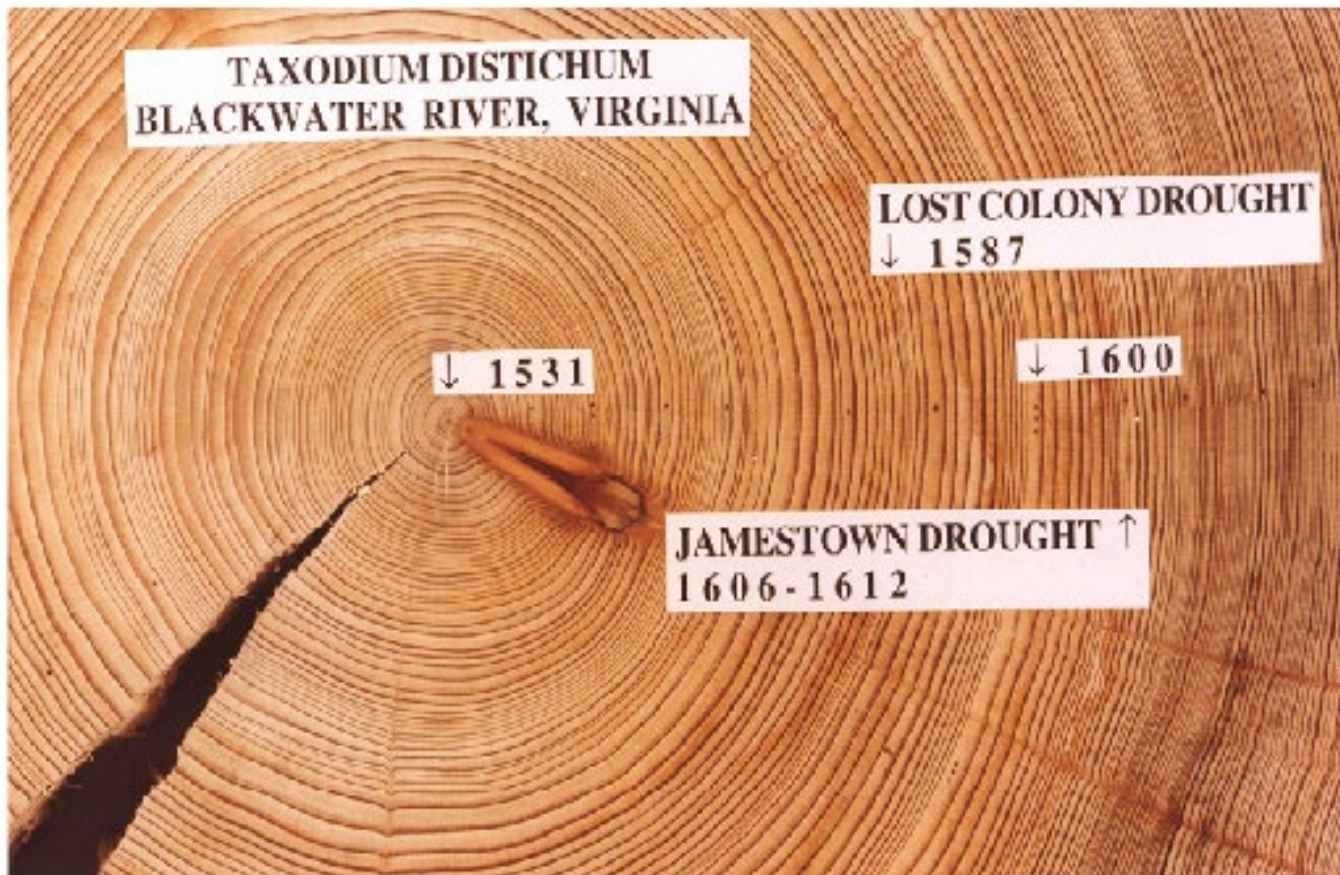
WHAT DO WE GET?

- Reconstructions can yield
 - Temperature time histories



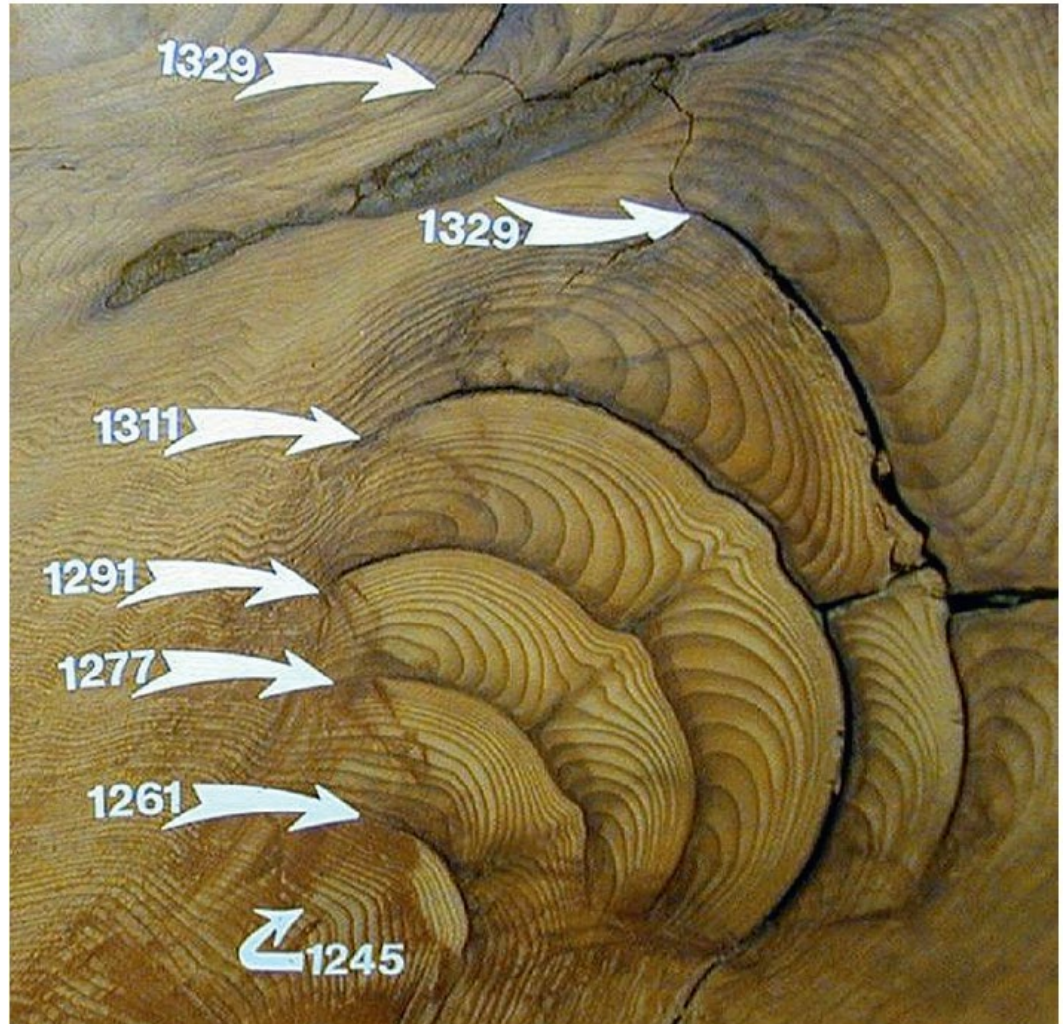
WHAT DO WE GET?

- Drought information:



WHAT DO WE GET?

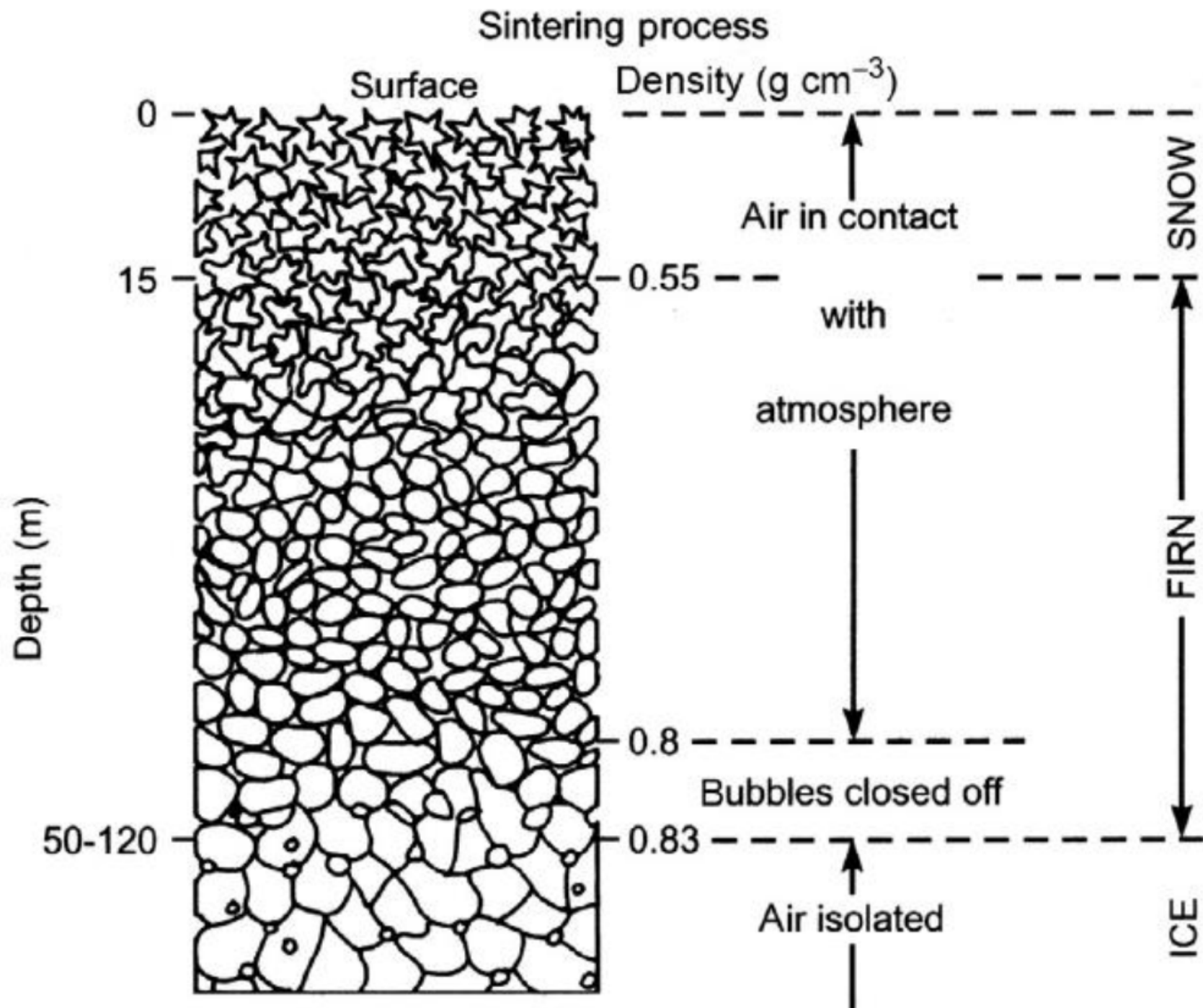
- Wildfire information.
- Which we can relate to other atmospheric variables.



ICE CORE BASICS

- Starts as snowfall.
- Snow accumulates ... slowly.
 - Interior regions of Antarctica receive only about 2 inches of snow a year.
- Weight of accumulated snow causes underlying crystals to settle, deform, and recrystallize, which removes the air spaces between the crystals.
 - Taking air out → increase in density
 - "Densification"









WHAT CAN WE GET FROM ICE CORES?

TABLE 5.1 Principal Sources of Paleoclimatic Information from Ice Cores

Parameter	Analysis
Temperature	
Summer	Melt layers
Annual	δD , $\delta^{18}O$ (ice), Ar, N ₂ (diffusion)
Source region temperature/humidity	Deuterium excess (<i>d</i>)
Accumulation (net)	Seasonal signals
Volcanic activity	Conductivity, non-sea salt (nss.) SO ₄ ²⁻ , glaciochemistry
Tropospheric turbidity	ECM, microparticle content, trace elements
Wind speed	Particle size, concentration
Atmospheric composition	Trace gases (CO ₂ , CH ₄ , and N ₂ O)
Sea ice extent	Glaciochemistry (Br ⁻ , I ⁻ , and Na ⁺)
Atmospheric circulation	Glaciochemistry (major ions)
Solar activity and geomagnetic field changes	¹⁰ Be
Forest fire history	Levoglucosan and other biomarkers



GEOLOGIC TIME SCALE



INTERNATIONAL CHRONOSTRATIGRAPHIC CHART

www.stratigraphy.org

International Commission on Stratigraphy













v 2018/08



Eonothem / Eon Erathem / Era System / Period	Series / Epoch	Stage / Age	GSSP	numerical age (Ma)
Phanerozoic	Quaternary	Holocene	UL	present
		Meghalayan	M	0.0002
		Northgrippian	L	0.0082
		Greenlandian	L	0.0117
		Pleistocene		0.126
		Middle		0.781
		Calabrian		1.80
		Gelasian		2.58
		Pliocene		3.600
		Zanclean		5.333
	Neogene	Messinian		7.246
		Tortonian		11.63
		Serravallian		13.82
		Langhian		15.97
		Burdigalian		20.44
		Aquitanian		23.03
		Chattian		27.82
		Oligocene		33.9
		Rupelian		37.8
		Priabonian		41.2
	Paleogene	Eocene		47.8
		Lutetian		56.0
		Ypresian		59.2
		Thanetian		61.6
		Selandian		66.0
		Danian		72.1 ± 0.2
		Maastrichtian		83.6 ± 0.2
		Campanian		86.3 ± 0.5
		Santonian		89.8 ± 0.3
		Coniacian		93.9
Mesozoic	Cretaceous	Turonian		100.5
		Cenomanian		~ 113.0
		Albian		~ 125.0
		Aptian		~ 129.4
		Barremian		~ 132.9
		Hauterivian		~ 139.8
		Valanginian		~ 145.0
		Berriasian		

Eonothem / Eon Erathem / Era System / Period		Series / Epoch	Stage / Age	GSSP	numerical age (Ma) ~ 145.0		
Phanerozoic	Mesozoic	Jurassic	Upper	Tithonian		152.1 ±0.9	
				Kimmeridgian		157.3 ±1.0	
			Middle	Oxfordian		163.5 ±1.0	
				Callovian		166.1 ±1.2	
				Bathonian	🚩	168.3 ±1.3	
		Bajocian		🚩	170.3 ±1.4		
		Aalenian		🚩	174.1 ±1.0		
		Lower	Toarcian	🚩			
			Pliensbachian	🚩	182.7 ±0.7		
			Sinemurian	🚩	190.8 ±1.0		
			Hettangian	🚩	199.3 ±0.3		
				🚩	201.3 ±0.2		
		Triassic	Upper	Rhaetian		~ 208.5	
				Norian		~ 227	
				Carnian	🚩	~ 237	
	Middle		Ladinian	🚩	~ 242		
			Anisian		247.2		
	Lower		Olenekian		251.2		
			Induan	🚩	251.902 ±0.024		
	Paleozoic		Permian	Lopingian	Changhsingian	🚩	254.14 ±0.07
					Wuchiapingian	🚩	259.1 ±0.5
				Guadalupian	Capitanian		265.1 ±0.4
		Wordian			🚩	268.8 ±0.5	
		Cisuralian		Roadian	🚩	272.95 ±0.11	
				Kungurian		283.5 ±0.6	
Artinskian					290.1 ±0.26		
Sakmarian				🚩	293.52 ±0.17		
Asselian				🚩	298.9 ±0.15		
Gzhelian					303.7 ±0.1		
Carboniferous	Pennsylvanian	Upper	Kasimovian		307.0 ±0.1		
			Moscovian		315.2 ±0.2		
	Middle						
		Lower	Bashkirian	🚩	323.2 ±0.4		
	Mississippian		Upper	Serpukhovian		330.9 ±0.2	
		Middle					
		Lower	Visean	🚩	346.7 ±0.4		
			Tournaisian	🚩	358.9 ±0.4		

Eonothem / Eon Erathem / Era System / Period		Series / Epoch	Stage / Age	GSSP	numerical age (Ma) 358.9 ± 0.4	
Phanerozoic	Paleozoic	Devonian	Upper	Famennian		372.2 ±1.6
				Frasnian		382.7 ±1.6
			Middle	Givetian		387.7 ±0.8
				Eifelian		393.3 ±1.2
			Lower	Emsian		407.6 ±2.6
				Pragian		410.8 ±2.8
		Silurian	Lochkovian		419.2 ±3.2	
			Pridoli		423.0 ±2.3	
			Ludlow	Ludfordian		425.6 ±0.9
			Wenlock	Gorstian		427.4 ±0.5
			Llandovery	Homerian		430.5 ±0.7
				Sheinwoodian		433.4 ±0.8
	Ordovician	Upper	Telychian		438.5 ±1.1	
			Aeronian		440.8 ±1.2	
			Rhuddanian		443.8 ±1.5	
			Hirnantian		445.2 ±1.4	
			Katian		453.0 ±0.7	
			Sandbian		458.4 ±0.9	
	Lower	Darriwilian		467.3 ±1.1		
		Dapingian		470.0 ±1.4		
		Floian		477.7 ±1.4		
		Tremadocian		485.4 ±1.9		
	Mesozoic	Cambrian	Furongian	Stage 10		~ 489.5
				Jiangshanian		~ 494
Paibian					~ 497	
Miaolingian			Guzhangian		~ 500.5	
			Drumian		~ 504.5	
			Wuliuan		~ 509	
Terreneuvian		Series 2	Stage 4		~ 514	
			Stage 3		~ 521	
		Stage 2		~ 529		
			Fortunian		541.0 ±1.0	

Eonothem / Eon	Erathem / Era	System / Period	GSSP	numerical age (Ma)	
Precambrian	Proterozoic	Ediacaran		541.0 ± 1.0	
		Neo-proterozoic	Cryogenian		~ 635
			Tonian		~ 720
		Meso-proterozoic	Stenian		1000
			Ectasian		1200
			Calymmian		1400
			Paleo-proterozoic	Statherian	
		Orosirian			1800
		Rhyacian			2050
		Siderian			2300
	Meso-archean		2800		
	Paleo-archean		3200		
	Eo-archean		3600		
			4000		
	Hadean				~ 4600

Units of all ranks are in the process of being defined by Global Boundary Stratotype Section and Points (GSSP) for their lower boundaries, including those of the Archean and Proterozoic, long defined by Global Standard Stratigraphic Ages (GSSA). Charts and detailed information on ratified GSSPs are available at the website <http://www.stratigraphy.org>. The URL to this chart is found below.

Numerical ages are subject to revision and do not define units in the Phanerozoic and the Ediacaran, only GSSPs do. For boundaries in the Phanerozoic without ratified GSSPs or without constrained numerical ages, an approximate numerical age (~) is provided.

Ratified Subseries/Subepochs are abbreviated as UL (Upper/Late), M (Middle) and LE (Lower/Early). Numerical ages for all systems except Quaternary, upper Paleogene, Cretaceous, Triassic, Permian and Precambrian are taken from 'A Geologic Time Scale 2012' by Gradstein et al. (2012), those for the Quaternary, upper Paleogene, Cretaceous, Triassic, Permian and Precambrian were provided by the relevant ICS subcommissions.

Colouring follows the Commission for the Geological Map of the World (<http://www.cgmw.org>)

Chart drafted by K.M. Cohen, D.A.T. Harper, P.L. Gibbard, J.-X. Fan (c) International Commission on Stratigraphy, August 2018

To cite: Cohen, K.M., Finney, S.C., Gibbard, P.L. & Fan, J.-X. (2013; updated) The ICS International Chronostratigraphic Chart. Episodes 36: 199-204.

URL: <http://www.stratigraphy.org/ICSchart/ChronostratChart2018-08.pdf>



EARTH'S CLIMATE HISTORY

- We are currently in the:
 - Phanerozoic Eon
 - The Cenozoic Era
 - Quaternary Period
 - Holocene Epoch
- The previous Epoch, the Pleistocene, was characterized by the Ice Age



EARTH'S CLIMATE HISTORY: THE ICE AGE

- Ice Age: A series (cycle) of glacial and interglacial periods that occurred during a ~1.7 million year period of the Pleistocene



- Climate was extremely variable. A major player on the Northern American Continent was the Laurentide Ice Sheet



18,000 YEARS AGO



ICE AGE

- Each glacial advance erases evidence of the previous advance.
- Thus, our most reliable evidence extends only back to the last glacial maximum (LGM)
- The LGM began about 27K ya and peaked ~18K ya.

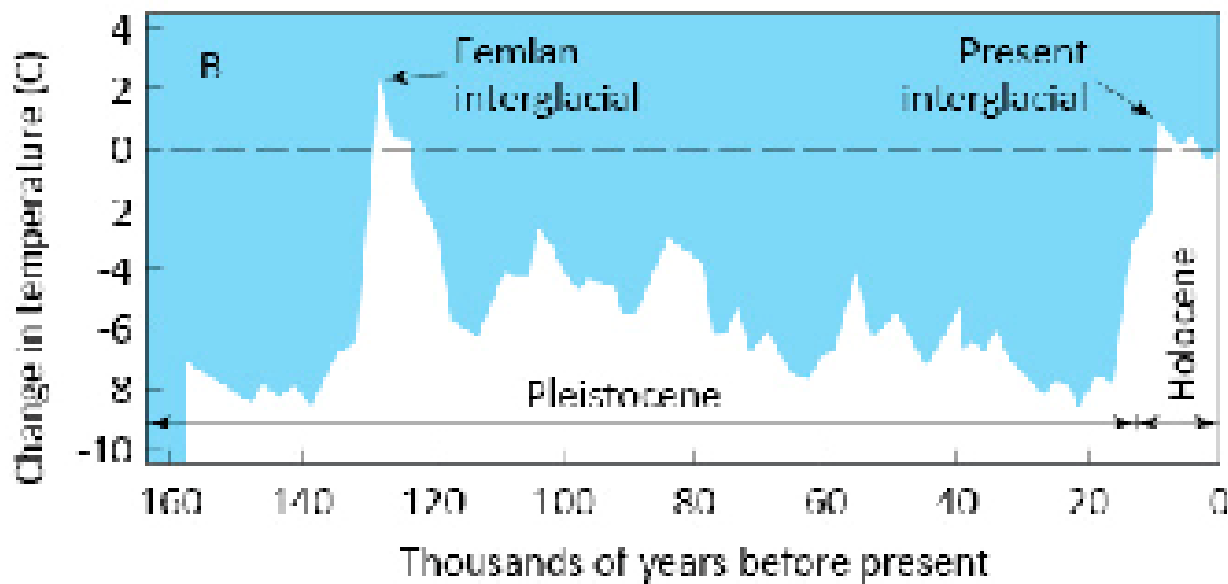
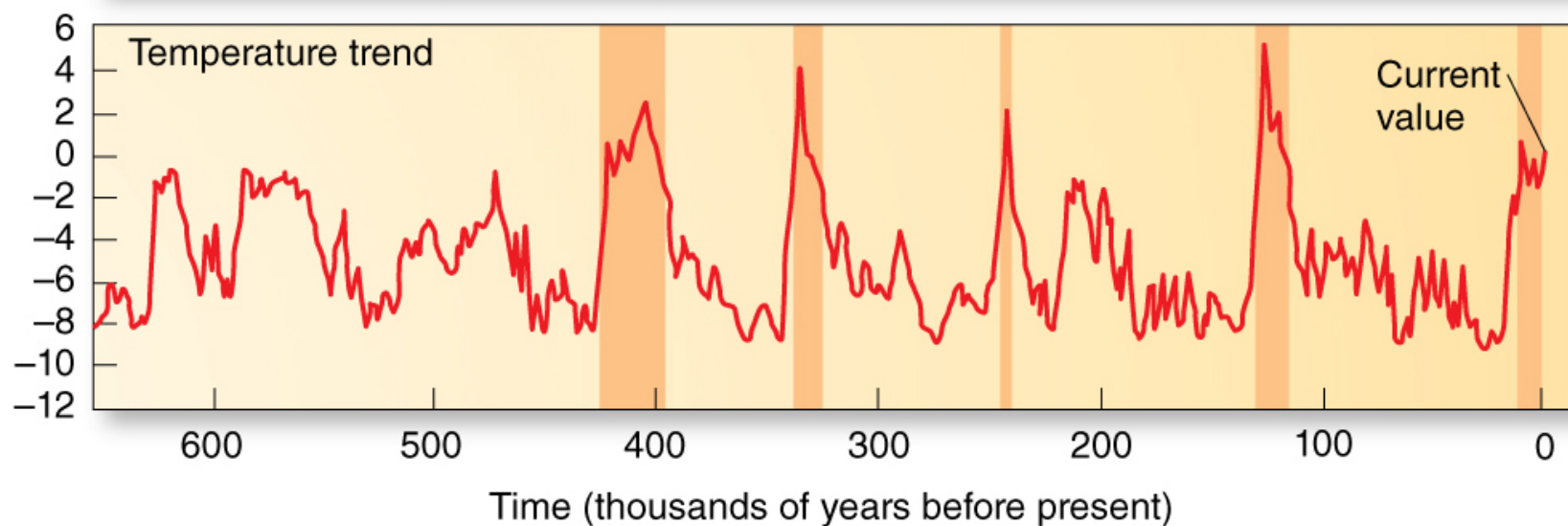


ICE AGE CYCLE

- The end of the LGM was marked by a temp increase to the present interglacial:
- Laurentide Ice Sheet began to melt.
- Melting occurred with only:
 - 5 Deg C temp increase in tropics
 - 6-8 Deg C temp increase in the mid-lats
 - 10 Deg C temp increase at the poles
- One of multiple cycles. Glacial period ~ every 100,000 years separated by ~11,000 year interglacial period.



Temperature change ($^{\circ}\text{C}$)



ICE AGE CYCLES

- Why were these period so repetitive?



- The 100,000 year fluctuation can be attributed to Milankovitch Theory: slight changes in earth's orbit alter the sun-earth geometry.

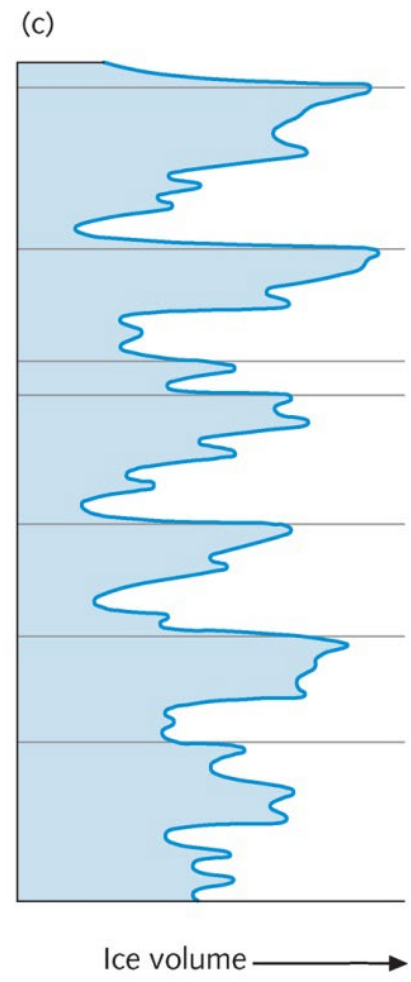
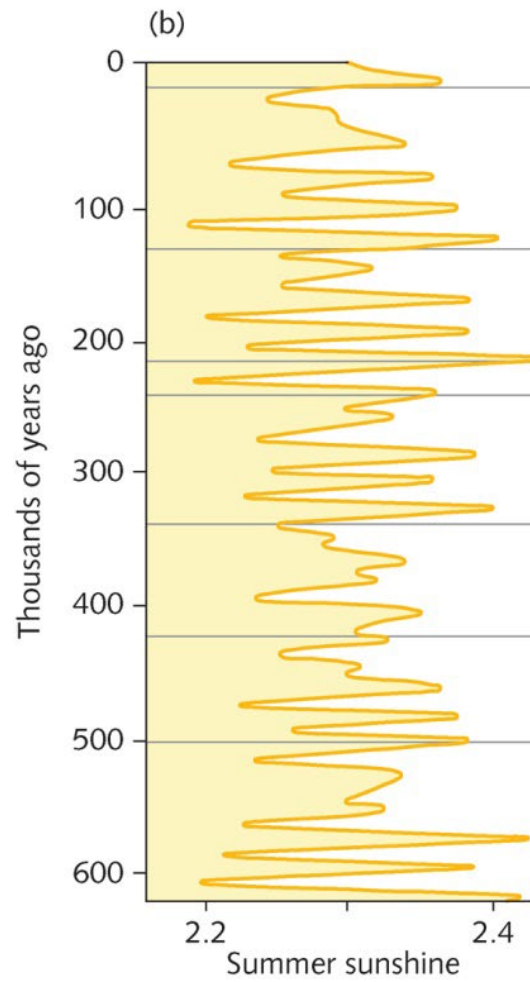
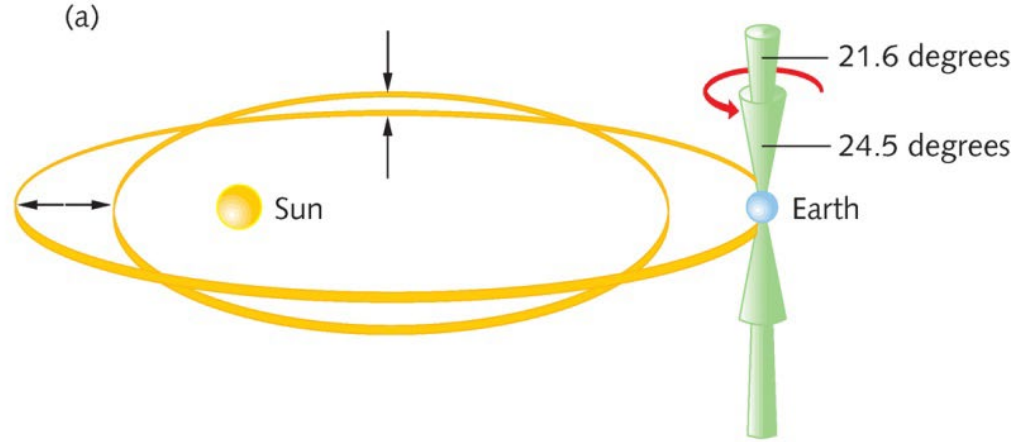


MILANKOVITCH THEORY

- Three modes:
 - Changes in the eccentricity of Earth's Orbit
 - How close to a perfect circle is Earth's orbit around the sun?
 - 100,000 year cycle
 - Changes in the obliquity of Earth's Orbit
 - Tracks changes in Earth's tilt
 - 41,000 year cycle
 - Precession
 - Which star is Earth's axis pointing to? Polaris right now.
 - 27,000 year cycle.



- Milankovitch modes explain about 60% of the variance in Earth's climate history.
- Effects greatest at high latitudes in summer where insolation can vary by 20%
- These changes (glacial vs interglacial) have been globally synchronous



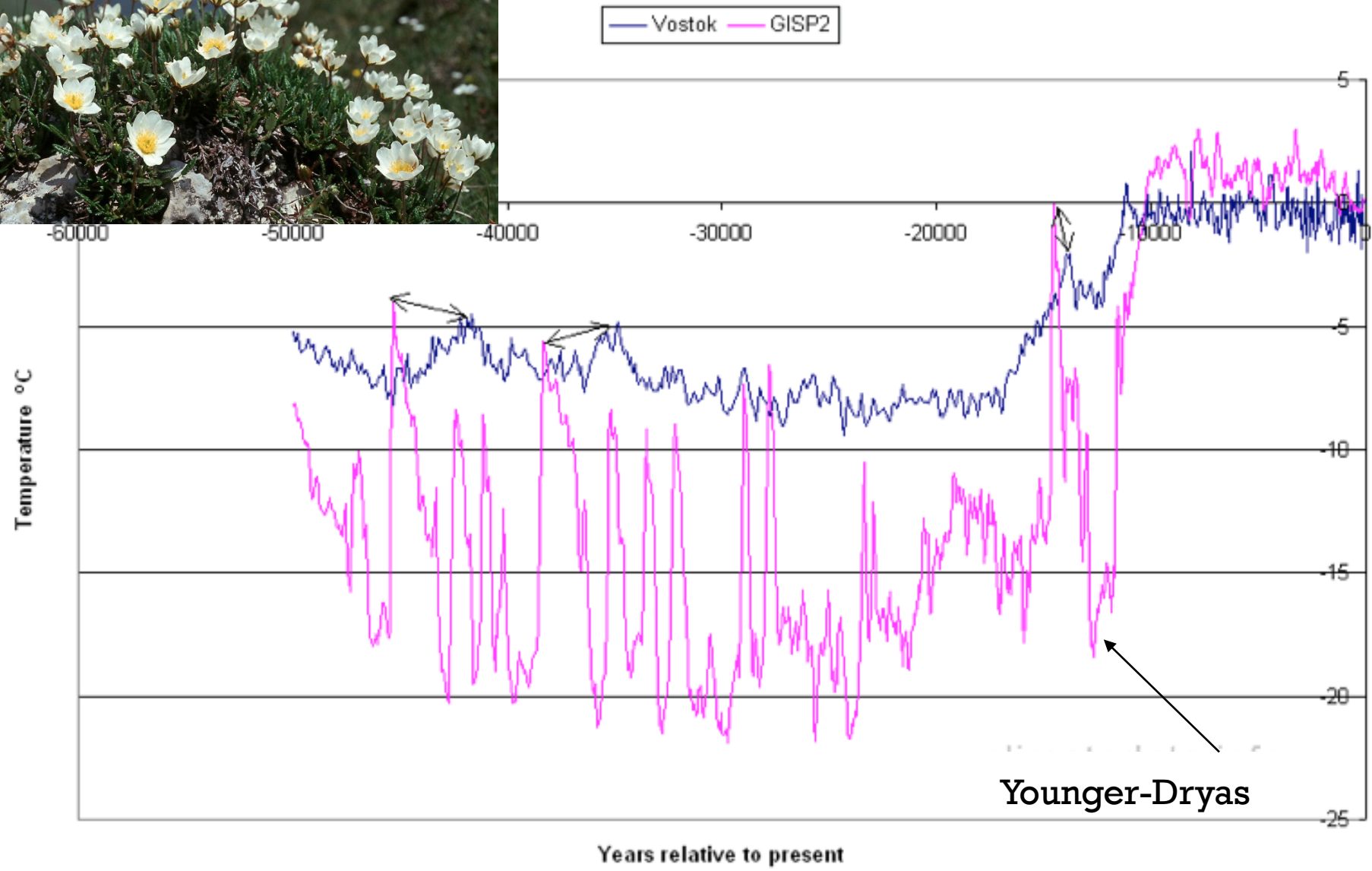
GLOBAL / NON-GLOBAL CHANGES

- Despite glacial / interglacial periods being globally synchronous, ice core data show:
 - Southern Hemisphere climate been more stable
 - Time delays in most recent glacial cycle
 - Abrupt changes on the scale of 2-3K years and 7-12 K years.





Temperature - GISP2 and Vostok



YOUNGER-DRYAS

- Regional short-term climate fluctuation.
- Abrupt cooling
- Palynology data revealed the pollen of the Dryas Octopetala much further south than normal.
- Time scale too short to be Milankovitch...so what caused it?



YOUNGER DRYAS

- Related to the melting of the Laurentide Ice Sheet and the AMOC.



- When the AMOC resumed its normal pattern, warming continued
 - The arctic warmed by 7 °C over the course of 50 years .

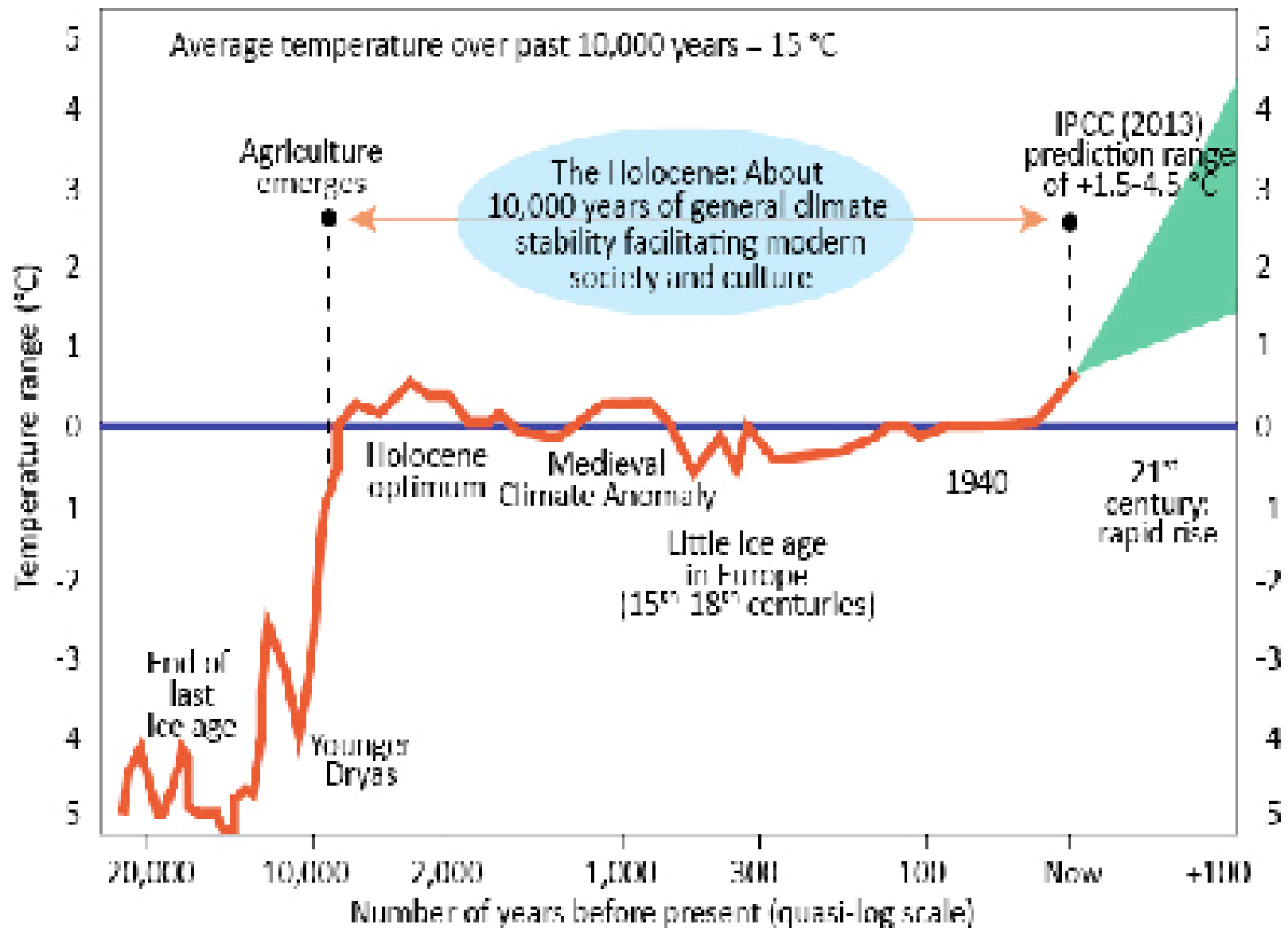


THE HOLOCENE

- At the conclusion of the Younger-Dryas Period, Earth entered into the present interglacial period: The Holocene Epoch (~10,500 YBP)
- Characterized by:
 - Relatively stable climate (compared to the last Epoch)
 - Development of Agrarian Societies
 - Modern Civilization



Temperature stability during the Holocene

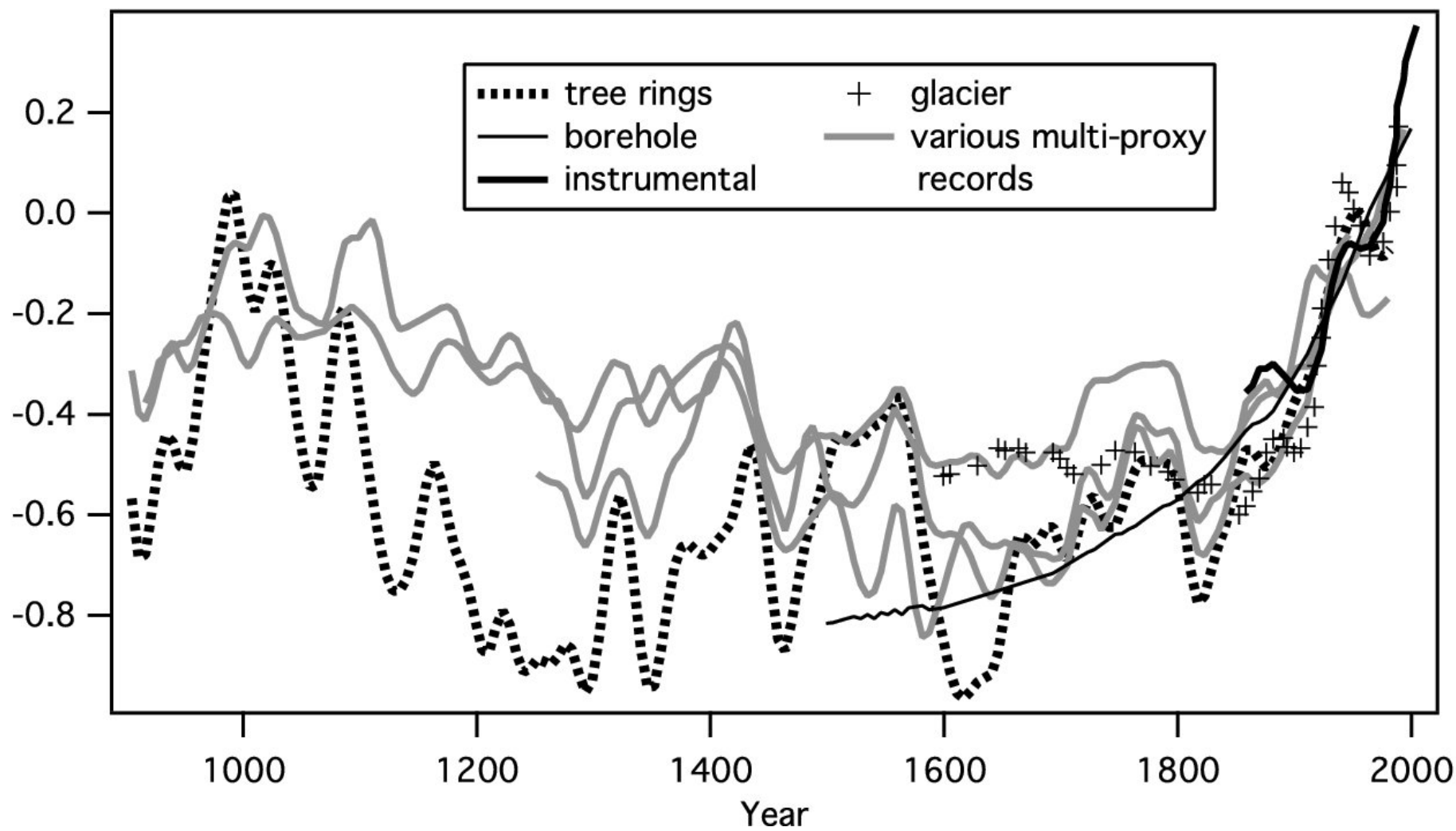


THE HOLOCENE

- There have also been several small scale climate fluctuations within the Holocene (specifically over the last 1000 years):
 - Medieval Climate Anomaly (MCA)
 - Mild Winters; Growing Wine in England
 - Isolated to N.H. based on data
 - Vikings get adventurous...
 - Little Ice Age (LIA)
 - Multi-century cool period
 - Global temp ~0.5 cooler.
 - Vikings give up; Food shortage; Societal Unrest.

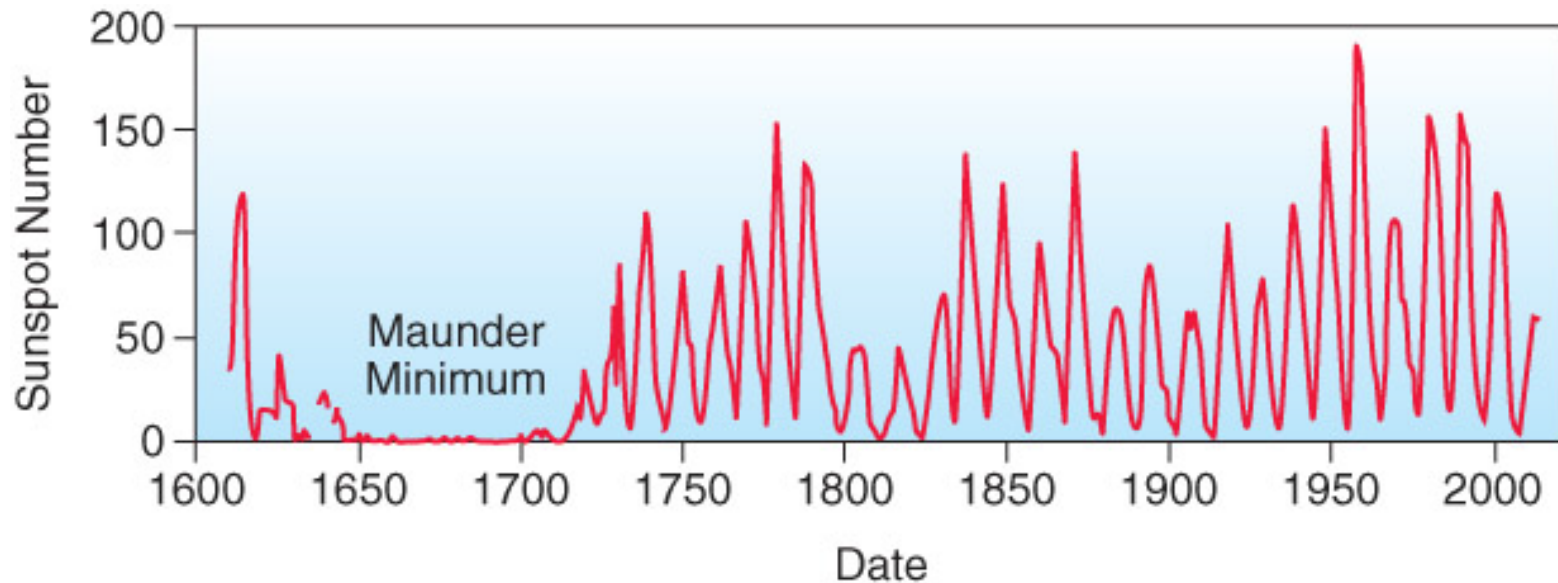


Temperature anomaly (°C)



LITTLE ICE AGE

- What caused it?
 - Not Entirely sure
 - Sunspots?



ANTHROPOCENE?

- Check out this article:

<https://www.smithsonianmag.com/science-nature/what-is-the-anthropocene-and-are-we-in-it-164801414/>

- Watch this clip:

<http://www.smithsonianmag.com/videos/category/science/what-is-the-anthropocene/?jwsourc=cl>



Global Warming 1850 to 2021

- HadCRUT5
- GISTEMP
- NOAA
- ECMWF
- Berkeley Earth

Global Temperature Anomaly ($^{\circ}\text{C}$)

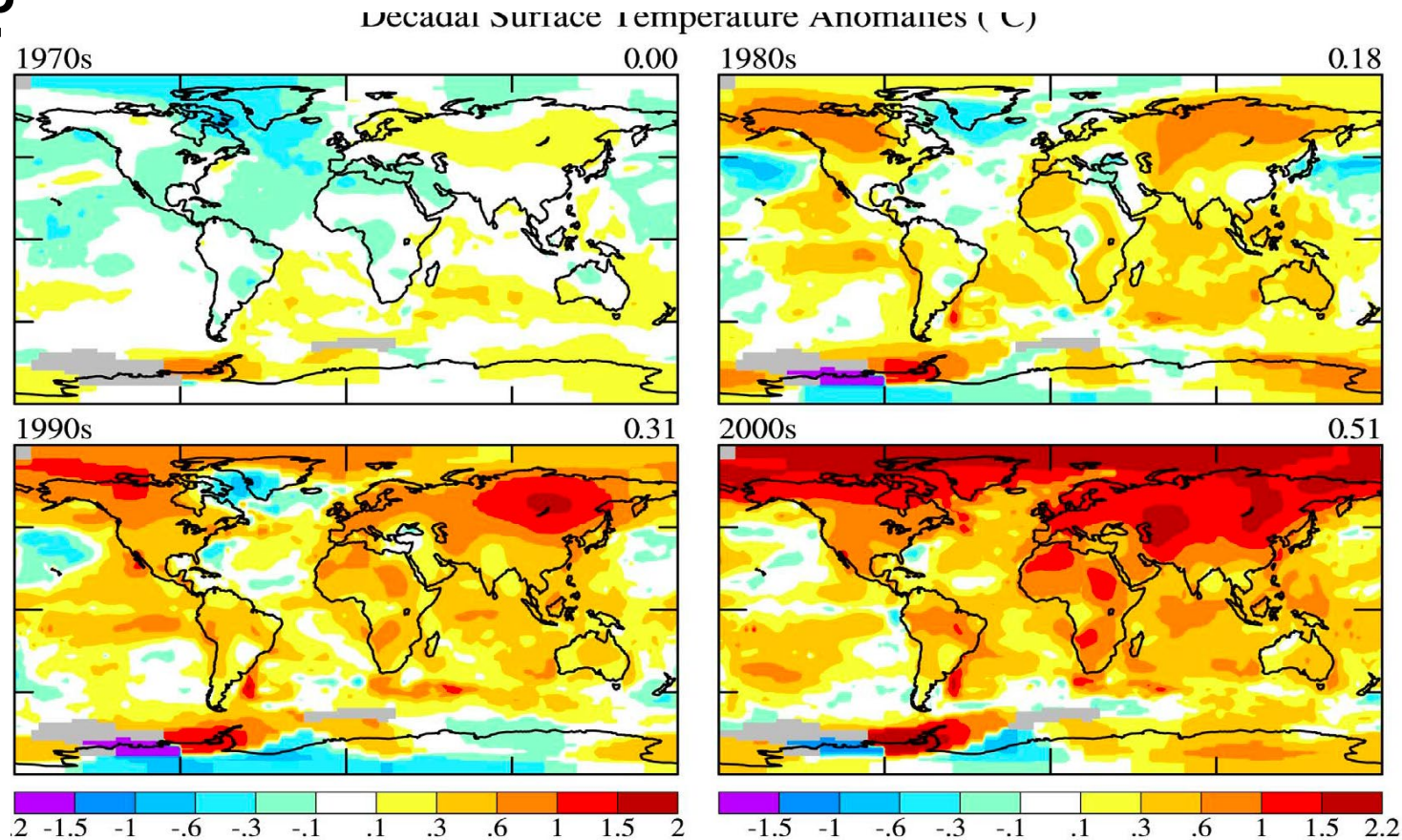
95% confidence interval shown for Berkeley Earth
Temperature anomalies relative to 1981-2010 average

1860 1880 1900 1920 1940 1960 1980 2000 2020



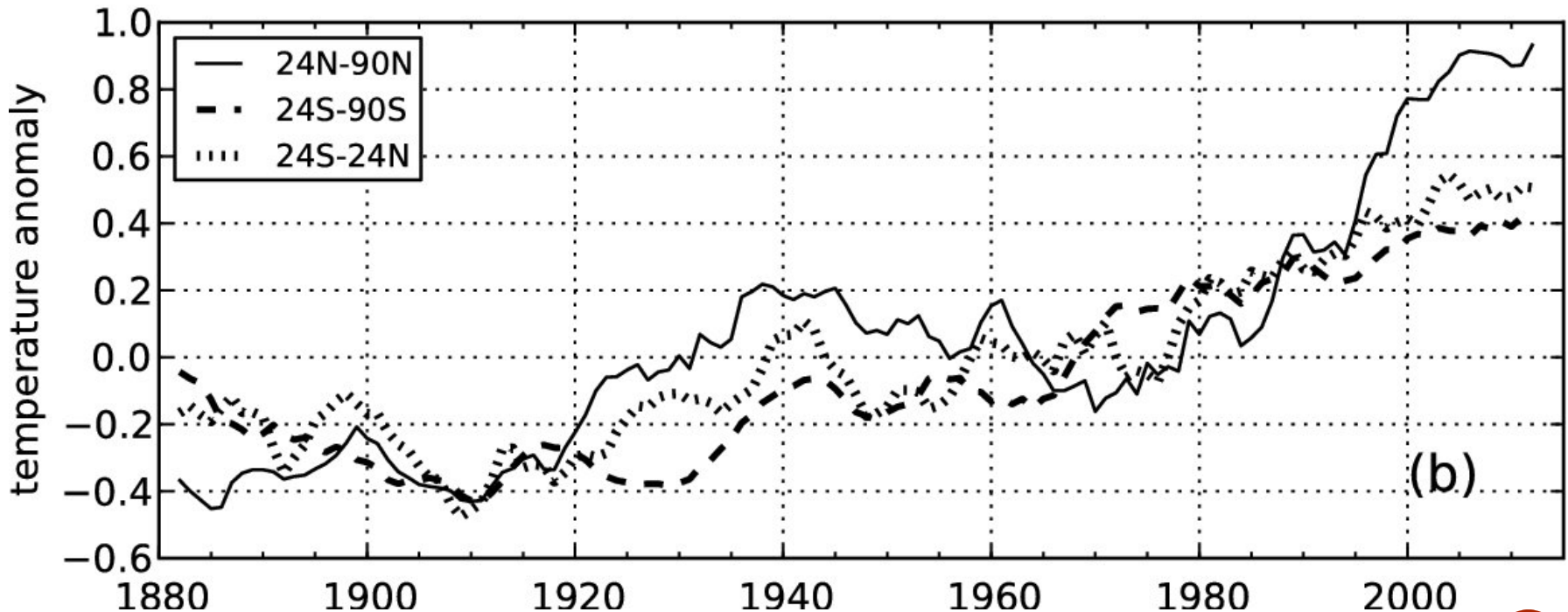
GLOBAL TEMPERATURE RECORD

- Additionally, land areas have warmed more than ocean
 - Why?



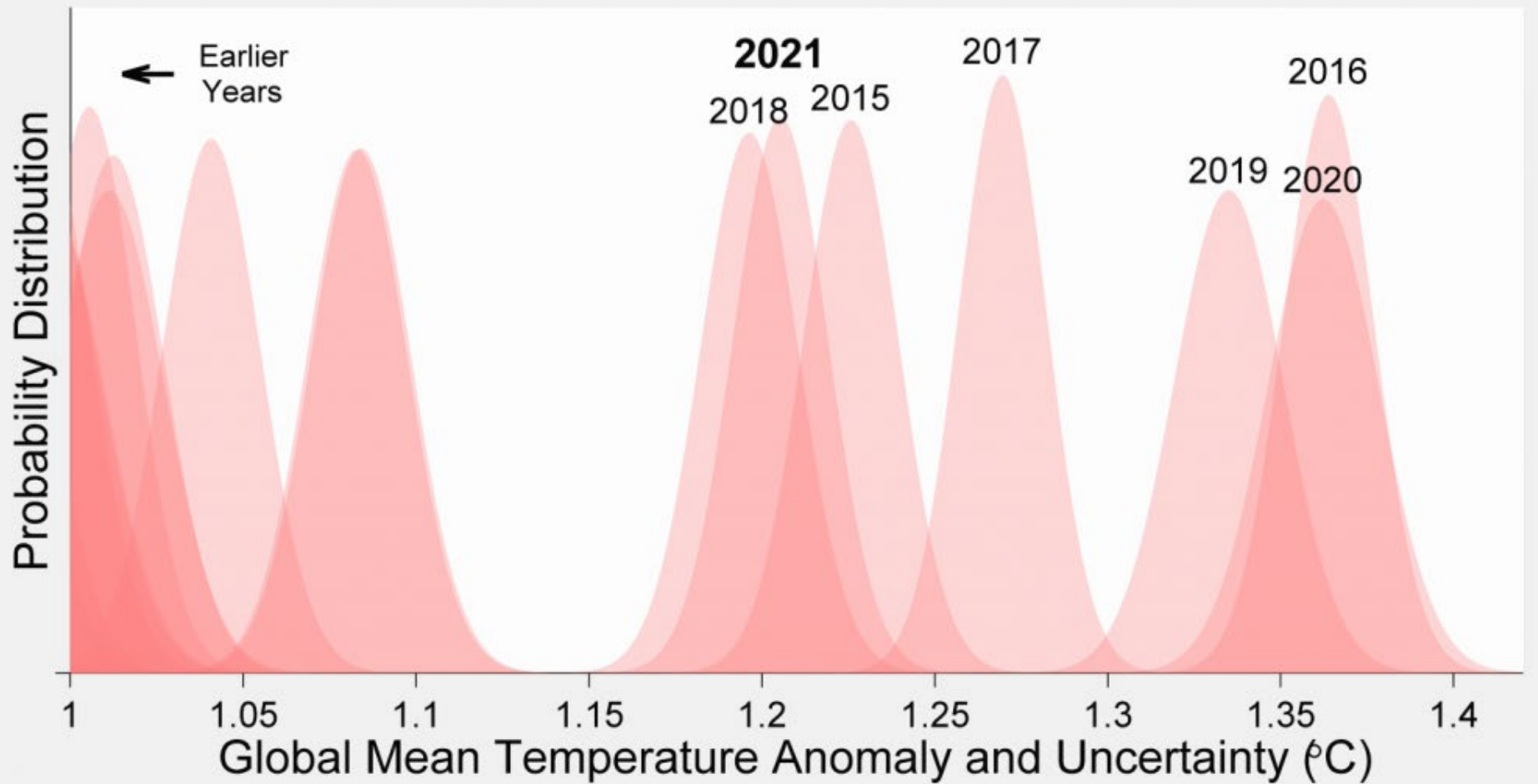
GLOBAL TEMPERATURE RECORD

- NH warmed more than Southern Hemisphere
 - Why?



RECENT DATA

Rank 1 = Warmest Period of Record: 1880–2021	Year	Anomaly °C	Anomaly °F
	2016	0.99	1.78
	2020	0.98	1.76
	2019	0.95	1.71
	2015	0.93	1.67
	2017	0.91	1.64
	2021	0.84	1.51
	2018	0.82	1.48
	2014	0.74	1.33
	2010	0.72	1.30
10 (tied)	2013	0.67	1.21
10 (tied)	2005	0.67	1.21



Based on Berkeley Earth's estimates of the global annual average temperature increase relative to 1850-1900.

